

**Supporting information for:  
New insights into the cooling of the oceanic lithosphere  
from surface-wave tomographic inferences**

**Franck Latallier<sup>1</sup>, Paula Koelemeijer<sup>1</sup>, Andrew Walker<sup>1</sup>, Alessia Maggi<sup>2</sup>,  
Sophie Lambotte<sup>2</sup>, Christophe Zaroli<sup>2</sup>**

<sup>1</sup>Department of Earth Sciences, University of Oxford, Oxford, United Kingdom

<sup>2</sup>Institut Terre et Environnement de Strasbourg, UMR7063, Université de Strasbourg, EOST/CNRS,  
67084, Strasbourg, France

**Supporting information**

This supporting information file contains details on our finite-frequency SOLA tomography method, as well as Table 1, Figures S1 to S5, and references to seismic networks.

## Finite-frequency surface-wave tomography using SOLA

Here, we detail the tomographic method and the resolution filtering process. Further details can be found in Latallerie et al. (2025).

### Model parameterisation and setup

We measure the Rayleigh-wave phase-delays for a set of source-receiver-frequency triplets. We note  $d_i$  the datum for triplet  $i$  ( $1 \leq i \leq N$ ). For each triplet, we compute the associated 3D finite-frequency sensitivity kernel  $K_i(\mathbf{x})$  using the theory (and codes) derived by Zhou et al. (2004) and references therein. These sensitivity kernels take value at every location  $\mathbf{x}$  in the 3D space  $\oplus$  and relate the data linearly to the  $\delta \ln V_{SV}$  variations  $m(\mathbf{x})$  in the 3D space:  $d_i = \iiint_{\mathbf{x} \in \oplus} K_i(\mathbf{x}) m(\mathbf{x}) d^3 \mathbf{x}$ . This is fundamentally different from previous surface-wave models such as SS2DPacific (Latallerie et al., 2022) that consist of 2D layers.

We parameterise the 3D space into a grid of 3D voxels, indexed with  $j$  ( $1 \leq j \leq M$ ), each one with a volume  $V_j$ . In the following, we use the Einstein's summation convention, unless explicitly stated otherwise. Using this parameterisation, the forward problem becomes:  $d_i = G_{ij} m_j$ , where  $G_{ij} = K_{ij} V_j$  (without summation), and  $K_{ij}$  is the value of the sensitivity kernel  $i$  interpolated in cell  $j$ ,  $V_j$  the volume of cell  $j$ , and  $m_j$  the velocity variations  $\delta \ln V_{SV}$  interpolated in cell  $j$ . When considering many data, these linear relations can be combined and written in matrix form:  $\mathbf{d} = \mathbf{G} \mathbf{m}$ .

It is important to note that the forward problem relates the data directly to the 3D  $\delta \ln V_{SV}$  variations through only one linear equation. In contrast, typically employed approaches in surface-wave tomography require two steps (e.g. in the SS2DPacific model also presented), either first one that relates the raw data to 1D  $\delta \ln V_{SV}$  structure with depth along a certain path, followed by another step between this 1D path-averaged  $\delta \ln V_{SV}$  structure and the lateral (2D) distribution of  $\delta \ln V_{SV}$  at a certain depth; or first one that relates the raw data to the lateral (2D) distribution of  $\delta \ln V_{SV}$  at a certain frequency (phase or group-velocity maps), followed by another step that relates lateral variations in  $\delta \ln V_{SV}$  at various frequencies to the 1D distribution of  $\delta \ln V_{SV}$  versus depth at each lateral location. Neither of these approaches is set up in a truly 3D framework, such that any resolution analysis in these cases is inherently 2D or 1D (and uncertainty propagation needs to account for the two-step nature of the process).

### SOLA inferences

SOLA inference is local, conducted independently for each model parameter of interest (Zaroli, 2016). Let us concentrate on estimating  $\delta \ln V_{SV}$  in voxel  $k$ :  $\tilde{m}^{(k)}$ . Here, we use tilde ( $\tilde{\phantom{x}}$ ) to indicate that this is an estimate and we use the index  $k$  instead of  $j$  for reasons that will become clear later. Additionally, we use the upper-script notation to emphasize the local nature of the reasoning. We search for a model estimate that is a linear combination of the data:  $\tilde{m}^{(k)} = \mathbf{G}^\dagger_i^{(k)} d_i$ , where  $\mathbf{G}^\dagger_i^{(k)}$  is the vector that contains the weights of the linear combination. How should we build  $\mathbf{G}^\dagger_i^{(k)}$ ? To decide on this, we need to investigate the role it plays. Replacing  $d_i$  in the last expression using the forward relationship  $d_i = G_{ij} m_j$ , we obtain  $\tilde{m}^{(k)} = G_i^\dagger_i^{(k)} G_{ij} m_j$ . This tells us that the model estimate for parameter  $k$  is an average over all true model parameters, with the averaging given by the vector  $\mathbf{A}^{(k)}$  with components  $A_j^{(k)} = G_i^\dagger_i^{(k)} G_{ij}$ . This is the resolving kernel for model parameter  $k$  (though in practice we also take the volume of the voxels into account). The use of two letters,  $j$  or  $k$ , to identify the grid voxels becomes clear here, as  $j$  identifies the native physical model parameters  $\delta \ln V_{SV}$ , while  $k$  identifies a given local average.

The shape of the resolving kernel is controlled by  $\mathbf{G}^\dagger^{(k)}$ . Based on this geometrical interpretation, we thus want to design  $\mathbf{G}^\dagger^{(k)}$  such that the resolving kernel is as focused as possible around the location of model parameter  $k$ . To achieve this, we design a target resolving kernel, and solve an optimisation problem based on the minimisation of the squared differences between the target kernel and the resolving kernel. Last, but not least, for a linear forward problem, we can express the estimated model uncertainty  $C_{\tilde{m}_k}$  as a function of the data uncertainties  $C_{d_i}$ :  $C_{\tilde{m}_k} = G_i^{\dagger(k)2} C_{d_i}^2$ . Thus, we realise that  $\mathbf{G}^\dagger^{(k)}$  also controls the propagation of data uncertainty into model uncertainty. The higher the resolution, the stronger the propagation of uncertainty as well, leading to a resolution-uncertainty trade-off. SOLA computes  $\mathbf{G}^{(k)}$  by optimising the balancing between the fit to the target resolution and the uncertainty propagation. Finally, from  $\mathbf{G}^{(k)}$ , we can compute the model estimate, its uncertainty, and the 3D resolving kernel (see Figure 10 in Latallerie et al. (2025)).

The process described above can be repeated for every single (or a chosen set of) model parameters, which can be visualised as a *tomographic model*. However, it is now clear that every single model parameter in this *tomographic model* is in fact a local average of the true model parameters (with some associated uncertainty and limited resolution). The advantage of SOLA is two-fold: firstly, we have direct control over the model resolution, as we push it towards some target resolution (for example one may prefer a broader, but well behaved, circular-shaped resolution over a higher, but very irregular resolution); secondly, we obtain the resolution (and uncertainty) by construction for each model parameter of interest. It is worth repeating that, since our finite-frequency framework is fundamentally 3D, the resolving kernels, obtained for every model parameter in the 3D space, are also functions of the 3D space themselves. This is a major advance compared to former implementations of SOLA with surface waves (including model SS2DPacific), and it allows us to discuss 3D resolution effects in surface-wave tomography, especially with depth.

### Tomographic filtering of predictions

Imagine we have predictions for the  $\delta \ln V_{SV}$  structure ( $m_j^{\text{pred}}$ ) for a set of voxels  $j$ . The physical meaning of the predictions and the tomographic model are very different. While the prediction provides physical parameters at absolute locations in space, the tomographic model consists of a set of local averages of these parameters. Therefore, any meaningful comparison between the tomographic model and the predictions should start by applying the local averaging (through the resolving kernels) to the predictions; a process commonly known as ‘tomographic filtering’. If the resolution for the tomographic model is available (a rare commodity though, especially in 3D for surface-waves), the process is relatively straightforward: compute  $\widetilde{m^{\text{pred}}}^{(k)} = G_i^{\dagger(k)} G_{ij} m_j^{\text{pred}}$ .  $\widetilde{m^{\text{pred}}}^{(k)}$  is the filtered prediction, which is now directly comparable to the tomographic model.

In this study, we show that tomographic filtering has a large impact on predictions of the thermal structure of the oceanic lithosphere based on the plate cooling (PC) and half-space cooling (HSC) models. We use the modified plate cooling (PC) model from Richards et al. (2018), denoted as  $T^{\text{PC}}(t, z)$ , where  $z$  is the depth and  $t$  is the age of the lithosphere. To link geographic locations to age  $t(\theta, \phi)$ , we use the age model of Seton et al. (2020), where  $\theta$  and  $\phi$  denote latitude and longitude. We can therefore compute predictions at every location in the 3D tomographic grid:  $m_j^{\text{PC}} = T^{\text{PC}}(t(\theta_j, \phi_j), z_j)$ , where  $\theta_j$ ,  $\phi_j$ , and  $z_j$  are the latitude, longitude and depth of the centre of voxel  $j$ . Subsequently, we compute the filtered predictions as  $\widetilde{m^{\text{PC}}}^{(k)} = G_i^{\dagger(k)} G_{ij} m_j^{\text{PC}}$ , which we can then compare to the model estimates from our 3D surface-wave tomography. We follow the same process for the HSC model, except that the

temperature field is given by the equations and constants from Stein and Stein (1992):  $T^{\text{HSC}}(t, z) = T_m \text{erf}(z/\sqrt{4\kappa t})$ , where  $T_m = 1350^\circ\text{C}$  is the mantle potential temperature,  $\kappa = k/(\rho_m C_p)$  is the thermal diffusivity, with  $k = 3.138 \text{ J s}^{-1} \text{ m}^{-1} \text{ K}^{-1}$  the thermal conductivity and  $C_p = 1171 \text{ J K}^{-1} \text{ kg}^{-1}$  the specific heat capacity, and  $\text{erf}(\cdot)$  the error function.

Using the above filtering approach, we demonstrate that the differences between the non-filtered and filtered predictions (namely between  $m^{\text{PC}}_j$  and  $\widehat{m^{\text{PC}}^{(k)}}$ , as well as between  $m^{\text{HSC}}_j$  and  $\widehat{m^{\text{HSC}}^{(k)}}$ ) are large and cannot be ignored. Resolution does not only add random noise in the model, it introduces artificial structures that can be extremely misleading. This is particularly true with depth, where the complex sensitivity of surface waves with depth leads to a poorly behaved 3D resolution.

### Note on the parameterisation

The parameterisation should be chosen as small as possible. Any discrepancy between the continuous sensitivity kernels and their discretised counterparts represents a theoretical error that should be added to the global uncertainty. Therefore, the smaller the voxels, the smaller this uncertainty contribution. Visual screening of the sensitivity kernels show that they retain their main properties. This has been investigated more in detail in Latallerie et al. (2025).

Additionally, the parameterisation should be small enough to honor the conceptual shape of the continuous target kernels. In this study, the parameterisation is relatively coarse compared to the size of the target kernels, which, additionally, are chosen relatively small. However, as long as we are satisfied with the discretised versions of the target kernels themselves, there is no drawback. Here, the projected target kernels resemble ‘Dirac delta’-kernels, consequently the SOLA method attempts to construct resolution kernels that are as small as possible, given the limitations imposed by the uncertainty and chosen tradeoff parameter.

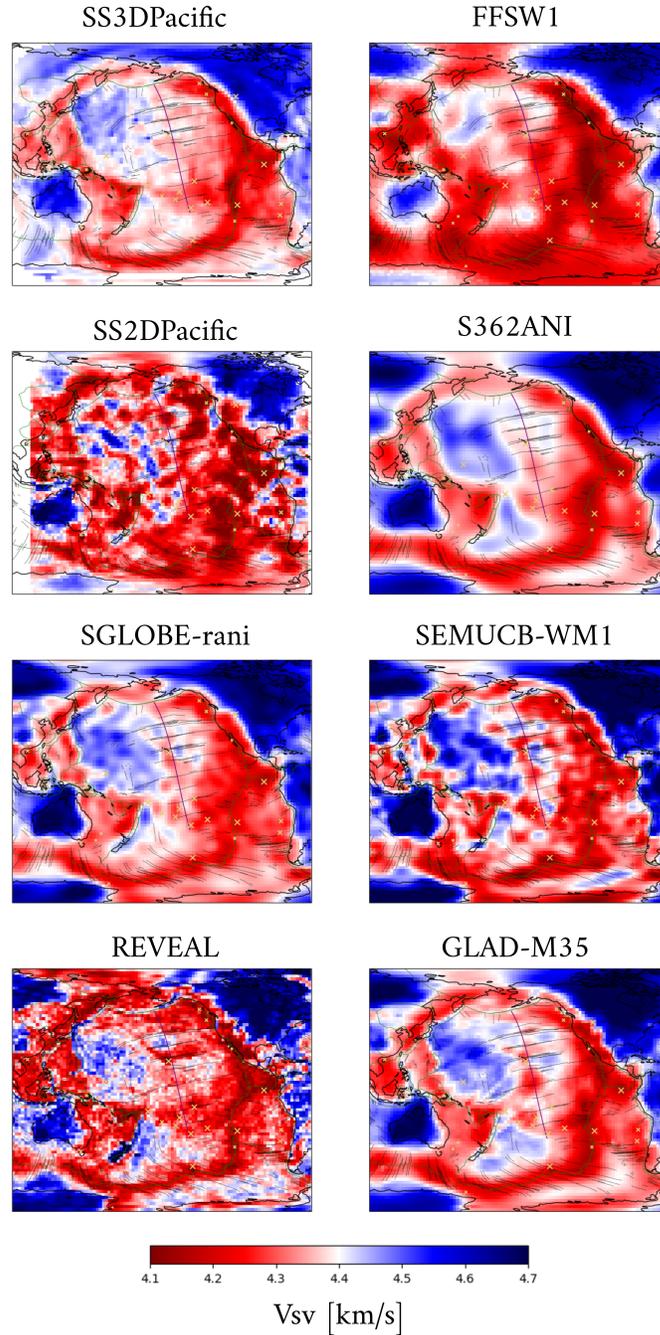
## Tables and figures

1

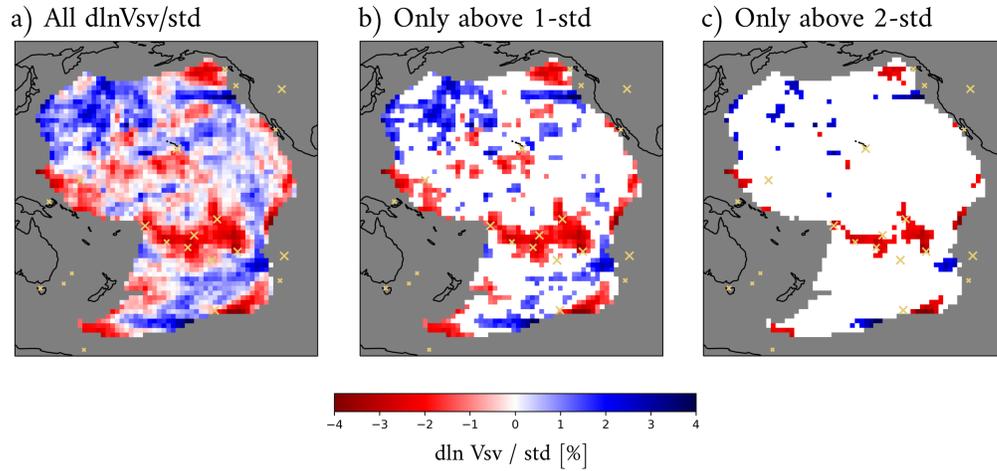
**Table 1.** Details of tomography models shown in Figure 5 and Figure S1

Name	Forward	Inverse	Parameters	Parameterisation	Data	Domain
SS3DPacific	Born FF	SOLA 3D	Vsv	Local	R fund. phase	Uppermost-mantle
SS2DPacific	Ray	Data-misfit + SOLA 2D	Vsv	Local	R multi. phase	Upper-mantle
FFSW1	Born FF	Data-misfit	Vsv, Vsh	Local	R/G fund. phase	Upper-mantle
REVEAL	Adjoint	Data-misfit	Vsv, Vsh, Vp	Local	Full-waveform	Whole-mantle
GLAD-M35	Adjoint	Data-misfit	Vsv, Vsh, Vp, eta	Local	Full-waveform	Whole-mantle
SEMUCB-WM1	SEM-NACT	Data-misfit	Vs, Vsh/Vsv	Local	Full-waveform	Whole-mantle
SGLOBE-RANI	Ray	Data-misfit	$V_s, \frac{V_{sh}^2 - V_{sv}^2}{2V_s^2}, \rho$	Global	R/G multi. phase/group, TT	Whole-mantle
S362ANI	Ray	Data-misfit	Vs, Vsh-Vsv, Vph-Vpv	Global	R/G fund. phase, body waveform, TT	Whole-mantle

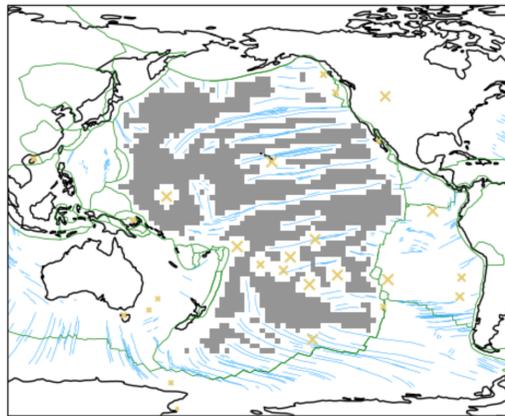
FF: Finite-frequency, R: Rayleigh, G: Love, fund.: fundamental, multi.: multimode, TT: Body-wave traveltimes,  $\rho$ : density.



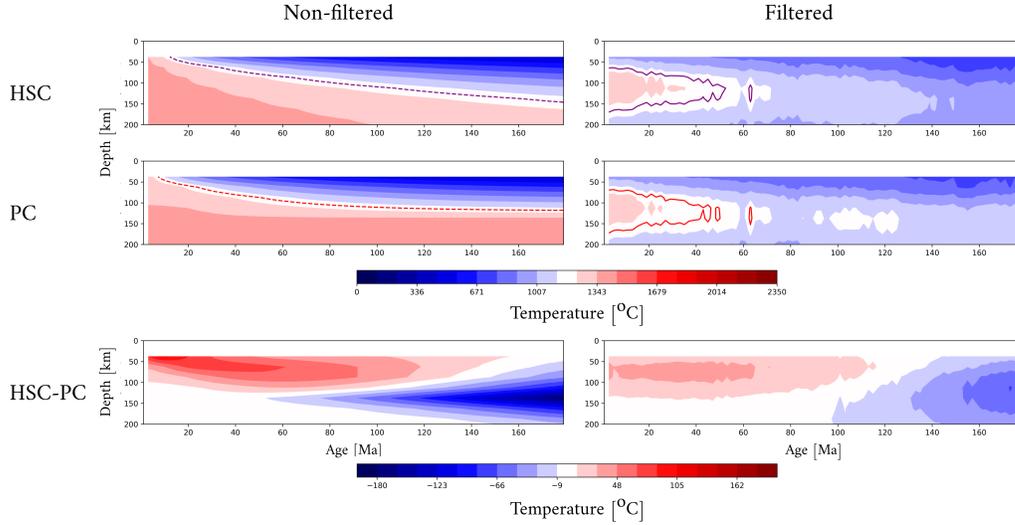
2 **Figure S1.** Comparison of the absolute  $V_{sv}$  structure of SS3DPacific compared to a selection of tomography models at 112 km depth. We include SS2DPacific (Latallerie et al., 2022),  
 3 REVEAL:(Thrastarson et al., 2024), FFSW1 (Zhou et al., 2006), S362ANI (Kustowski et al., 2008), SEMUCB-WM1 (French & Romanowicz, 2014), and  
 4 GLAD-M35 (Cui et al., 2024).  
 5  
 6



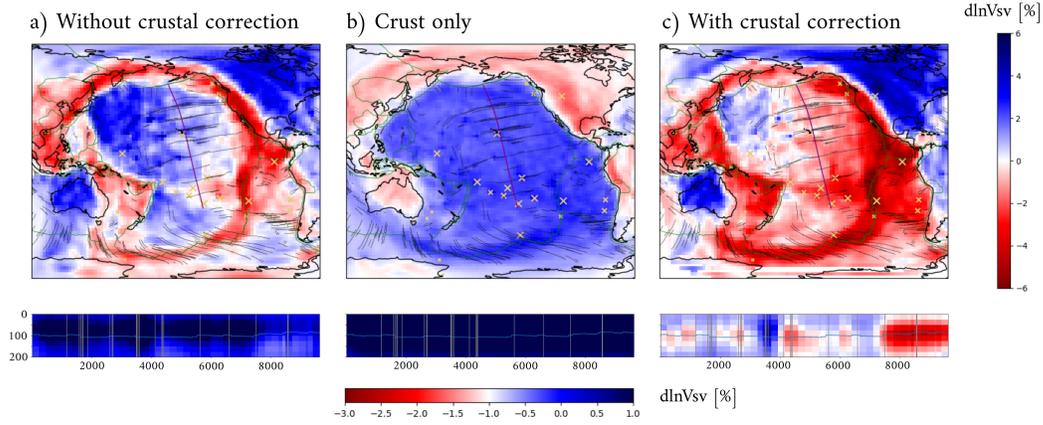
7 **Figure S2.** Significant lateral features in SS3DPacific, obtained by removing the age-  
 8 dependence trend (of SS3DPacific itself) and normalising by the model uncertainty (std). We  
 9 include a) all features, or mask values below b) the 1- $\sigma$  or c) the 2- $\sigma$  level. Yellow crosses indi-  
 10 cate hotspot locations, see the caption of Figure 1 in the main text.



11 **Figure S3.** Mask used to calculate the trends of the average velocity with age. Grey cells  
 12 indicate regions that are included, while regions near hotspots (yellow crosses), fracture zones  
 13 (black lines), and plate boundaries (green lines) are excluded. See the caption of Figure 1 in the  
 14 main text for more details.



15 **Figure S4.** Thermal structure according to models HSC (first row) and PC (second row),  
 16 as well as their differences (last row). The left column shows the original, non-filtered thermal  
 17 structure and the second column shows the filtered version using the SS3DPacific tomographic  
 18 filter. Purple and red dashed lines represent the 1175°C isotherm of the HSC and PC models, re-  
 19 spectively, while their solid counterparts represent the same isotherm, but for the filtered versions  
 20 of these models.



21 **Figure S5.** Assessment of influence of crustal structure on the model solution, showing results  
 22 at 112 km depth for SOLA inference of a) data without crustal corrections applied, b) only pre-  
 23 dictions for crustal structure, and c) data with crustal corrections (i.e. SS3DPacific). We use a  
 24 different colorscale for the cross-sections to better represent the range of velocities shown.

## Citations for seismic networks used in this study

We list here the network codes (in bold) for stations analysed in this study followed by the full citation as per the FDSN website, where available.

- **IP**: Anya Reading and Nick Rawlinson (2011); Ritter et al. (2014); Douglas Wiens and Maria Beatrice Magnani (2018); Sarah Mader et al. (2022)
- **6E** James Conder (2013); John Louie (2018); Pieter-Ewald Share and Frank Vernon (2019); Elnaz Seylabi et al. (2021); Dahm et al. (2023); Susan Bilek (2024)
- **7A** Ryberg and Haberland (2008); Ramon Carbonell (2012); Maureen Long and Paul Wiita (2013); Nicholas Schmerr (2017); Christopher W Johnson et al. (2018); Yuan et al. (2019); Marcelo Rocha (2021); Neil Harbison and Craig O'Neill (2024)
- **7C** Vergne et al. (2014); Rufus Catchings (2014); Derek Schutt and Rick Aster (2015); Costanzo et al. (2022)
- **7D** Brian Kennett (1997); Wes Thelen and Paul Bodin (2009); IRIS OBSIP (2011); Wenyuan Fan et al. (2018); Calum Chamberlain (2022)
- **AE** Arizona Geological Survey (2007)
- **AI** Istituto Nazionale di Oceanografia e di Geofisica Sperimentale (1992)
- **AK** Alaska Earthquake Center, Univ. of Alaska Fairbanks (1987)
- **AR** No DOI
- **AT** NOAA National Oceanic and Atmospheric Administration (USA) (1967)
- **AU** Geoscience Australia (2021)
- **AV** Alaska Volcano Observatory-USGS (1988)
- **AZ** Frank Vernon (1982)
- **BK** Northern California Earthquake Data Center (2014)
- **C** No DOI
- **C1** Universidad de Chile (2012)
- **CB** Institute of Geophysics China Earthquake Administration (IGPCEA) (2000)
- **CC** Cascades Volcano Observatory/USGS (2001)
- **CI** California Institute of Technology and United States Geological Survey Pasadena (1926)
- **CM** Servicio Geológico Colombiano (1993)
- **CN** Natural Resources Canada (1975)
- **CU** Albuquerque Seismological Laboratory (ASL)/USGS (2006)
- **DK** GEUS Geological Survey of Denmark and Greenland (1976)
- **G** Institut de physique du globe de Paris (IPGP) and École et Observatoire des Sciences de la Terre de Strasbourg (EOST) (1982)
- **GG** IRIS HQ (DC) (2012)
- **GS** Albuquerque Seismological Laboratory (ASL)/USGS (1980)
- **GT** Albuquerque Seismological Laboratory (ASL)/USGS (1993)
- **IC** Albuquerque Seismological Laboratory (ASL)/USGS (1992)
- **II** Scripps Institution of Oceanography (1986)
- **IM** Various Institutions (1965)
- **IU** Albuquerque Seismological Laboratory/USGS (1988)
- **JP** No DOI
- **KS** No DOI
- **LB** No DOI
- **LD** Lamont Doherty Earth Observatory (LDEO), Columbia University (1970)
- **MI** USGS Alaska Anchorage (2000)
- **MM** Department of Meteorology and Hydrology - National Earthquake Data Center (2016)

- **MX** Servicio Sismológico Nacional, Instituto de Geofísica, Universidad Nacional Autónoma de México, México (2017)
- **MY** No DOI
- **N4** Albuquerque Seismological Laboratory/USGS (2013)
- **NA** KNMI (2006)
- **NE** Albuquerque Seismological Laboratory (ASL)/USGS (1994)
- **NM** No DOI
- **NN** University of Nevada, Reno (1971)
- **NR** Utrecht University (UU Netherlands) (1983)
- **NU** Instituto Nicaraguense de Estudios Territoriales (INETER) (1975)
- **OO** NSF Ocean Observatories Initiative (2013)
- **PB** No DOI
- **PE** Penn State University (2004)
- **PN** No DOI
- **PR** University of Puerto Rico (1986)
- **PS** No DOI
- **PY** Frank Vernon (2014)
- **RM** Regional Integrated Multi-Hazard Early Warning System (RIMES Thailand) (2008)
- **RV** Alberta Geological Survey / Alberta Energy Regulator (2013)
- **S1** Michelle Salmon et al. (2011)
- **SC** New Mexico Tech (1999)
- **SV** No DOI
- **TA** IRIS Transportable Array (2003)
- **TM** No DOI
- **TW** Institute of Earth Sciences, Academia Sinica, Taiwan (1996)
- **TX** Bureau of Economic Geology, The University of Texas at Austin (2016)
- **UO** University of Oregon (1990)
- **US** Albuquerque Seismological Laboratory (ASL)/USGS (1990)
- **UW** University of Washington (1963)
- **WI** Institut De Physique Du Globe De Paris (IPGP) (2008)
- **XF** Frank Vernon and Jon Fletcher (1987); Steve Day and Frank Vernon (1994); Sridhar Anandkrishnan (1995); Frank Vernon and Ken Dueker (2000); John Nabelek (2002); Steve Grand and Jim Ni (2006); Chris Larsen and Michael West (2009); Douglas Wiens (2012); Meredith Nettles (2014); S. De Angelis and A. Diaz-Moreno (2017); Tiberi et al. (2020); Haberland et al. (2021); Anne Meltzer et al. (2021); Pilz et al. (2023); Paul et al. (2024)
- **XZ** Rick Aster and Philip Kyle (1999); Anthony Qamar (2001); Sylvie Leroy (2003); Roger Hansen and Gary Pavlis (2005); Chris Hayward (2014); Christian Poppeliers (2015); Frank Kruger et al. (2016); Silvio De Angelis (2018); Andy Nyblade (2019); SOCQUET et al. (2020)
- **Y6** Ian Joughin et al. (2006); Keir et al. (2016); Jamie Farrell et al. (2018); John Louie and Ronald Breitmeyer (2019); Tilmann et al. (2021); Weisen Shen (2023)
- **YP** Steve Grand and Jim Ni (2009); Zhao et al. (2016); Cornou et al. (2014); Gordon Hamilton (2016); Doubre et al. (2021); Amanda Bustin (2019)
- **Z7** Dan Klinglesmith (2008); Robert White (2010); Paul Bedrosian et al. (2017); Vergne and RESIF (2021); Tulane University et al. (2019); Dietrich Lange and Marcos Moreno (2023); Noah Finnegan and Susan Schwartz (2023)

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