

Status and challenges for monitoring ice motion and deformation with satellite constellations

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Why are we interested in polar ice masses ?



Polar ice masses are a key element in climate change and contribute to sea level rise.

Precise measurements of ice masses and velocity define boundary values for ice dynamic models and are needed for improving **forecasts**.

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- Methods for observing ice dynamics from SAR
- Standard ice velocity products for polar regions
- 3D ice velocity
- Mass change estimation from InSAR DEM differencing
- Summary for ice monitoring with SAR
- Upcoming missions



SAR Offset Tracking for Ice Motion

Matching of image templates by cross correlation

Amplitude correlation / speckle tracking / coherence tracking





- « Simultaneous » ascending and descending passes:
 - 4 displacement components to solve 3-D motion



Polar Ice Velocity Monitoring using Sentinel-1





Polar Ice Velocity Monitoring using Sentinel-1





Sentinel-1 Greenland Annual Ice Sheet Velocity Maps



Monthly Ice Velocity Products - since October 2014







Sentinel-1 Velocity Time series Jakobshavn Isbrae – West Greenland



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Sentinel-1 Velocity and Ice Flux Time Series Nunatakassaap Sermia – West Greenland



Ice velocity with InSAR

SAR interferometry measures motion along the LOS:

$$\Delta \phi = \frac{4\pi}{\lambda} \Delta R$$

3D velocity vector:
$$\mathbf{v} = \mathbf{v}_{h} + \mathbf{v}_{z} = v_{E} \mathbf{i}_{E} + v_{N} \mathbf{i}_{N} + v_{z} \mathbf{i}_{z}$$

Surface parallel flow assumption:

$$\mathbf{v} = \mathbf{v}_h + \mathbf{v}_z = \mathbf{v}_h + \mathbf{v}_x \frac{\partial}{\partial x} S(x, y) + \mathbf{v}_y \frac{\partial}{\partial y} S(x, y)$$

Vertical motion only due to slope

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Combination of ascending, descending interferograms and projection on DEM for retrieving 3-D **surface** velocity



Ice velocity with InSAR

ERS-1/-2 Interferogram 1995-10-31/1995-11-01

Drygalski Glacier, Antarctic Peninsula

Velocity Map (magnitude) from InSAR data of crossing orbits



3D Ice Motion with Offset Tracking



3D ice motion

TDX Stripmap mode:

- 3m resolution
- scene size (width x length) : 30 km x 50 km





Horizontal and vertical displacement rates



Intercomparison of Surface Elevation Changes



Geodetic mass balance from InSAR DEM differencing



Bistatic mode allows **elimination of temporal decorrelation** and atmospheric variations



- 1 (active) satellite is transmitting
- Both satellites receive simultaneously the radar echoes

Database for surface elevation change (SEC):

- TanDEM-X (TDM) bistatic InSAR time series (>3 years time steps)
- SRTM (where available)

Total net glacier mass balance from SEC:

$$\dot{B} = \rho \int_{A} \frac{\Delta h}{\Delta t} \, dA$$

 ρ – ice density

A - glacier basin area





DEM differencing and glacier mass change Antarctic peninsula

Analysis of Surface Elevation Change (dh/dt) 2011 - 2016

- Use of TanDEM-X data acquired from same satellite track in winter (frozen snow/ice) \rightarrow data sets with same σ° and same incidence angle
- Comparison TanDEM-X dh/dt with airborne lidar (ATM) data of NASA IceBridge, both data sets independently processed





DEM differencing and glacier mass change Antarctic peninsula



Sentinel-1 Brunt Ice Shelf Rift Monitoring



TCD, 2019, J. de Rydt et al.

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Main limitations



Upcoming SAR missions

Copernicus mission:

ROSE-L (Sentinel Expansion)

ESA EE10 Candidate missions:

Hydroterra (geostationnary C-band SAR)

HARMONY (SP-InSAR with S1)

NASA mission:

NISAR (L-band left-looking SAR)

https://cryoportal.enveo.at

opernicus